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Evidence of non-universality of von Kármán's κ

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Content

1. Introduction

2. Flows with Low Submergence

3. Flows with Bed-Load Transport

4. Flows with Suspended-Load Transport

5. Conclusions

Introduction

- ✓ The *von Kármán constant* κ relates the $u(y)$ profile along y in a wall-bounded shear flow to the **shear stress** at the bed surface.
- ✓ The universal logarithmic law of time-averaged streamwise velocity by Theodore von Kármán in 1930 is:

$$\frac{u(y)}{u_\tau} = \frac{1}{\kappa} \ln \left(\frac{y}{y_0} \right)$$

u_τ = shear velocity;

y_0 = zero-velocity level;

κ = von Kármán's constant, having a universal value of 0.41.

- ✓ The zero-velocity level y_0 is dependent on **smooth**, **transitional** and **rough** flow regimes governed by the shear Reynolds number R_* and/or k_s ;

$$R_* = u_\tau k_s / \nu$$

k_s = Nikuradse equivalent sand roughness;
 ν = kinematic viscosity of fluid.

- ✓ The expressions of y_0 for different flow regimes are (van Rijn 1993):

$$y_0(R_* \leq 3) = \frac{\nu}{9.1u_\tau} \quad \text{for smooth regime}$$

$$y_0(3 < R_* \leq 70) = \frac{\nu}{9.1u_\tau} + \frac{k_s}{30} \quad \text{for transitional regime}$$

$$y_0(R_* > 70) = \frac{k_s}{30} \quad \text{for rough regime}$$

- ✓ There have been some attempts to compute κ **mathematically** for an idealized flow over a **smooth rigid wall** or for homogeneous turbulence. For instance:
- 1) for an idealized flow, Long et al. (1993) analyzed the turbulent flow in **smooth pipes** with an eddy viscosity closure, obtaining $\kappa = 0.408 \pm 0.004$;
 - 2) Lo et al. (2005) estimated von Kármán's κ for wall-bounded turbulence based on the similarity with homogeneous constant-shear turbulence, obtaining $\kappa \approx 0.42$;
 - 3) using an energy argument and the mathematical symmetry for an idealized flow, Hughes (2007) argued that $\kappa = 0.414$ for **smooth wall**.

- ✓ Some considerations were made to set the u -profiles over rough beds and for low submergence by fixing the y_0 value.
- ✓ Koll (2006) described the parameterization of the u -profiles over rough beds by identifying three layers:
 - 1) the roughness layer, in which the u -profile is influenced by wall roughness;
 - 2) the log-law layer, in which the log-law velocity profile exists;
 - 3) the outer layer, which covers up to the free surface.
- ✓ The first two layers form the wall region.

- ✓ The u -profile holds, in the log-law layer, as

$$u/u_\tau = \kappa^{-1} \ln[(y - d)/(y_R - d)] + (u_R/u_\tau),$$

d = zero-plane displacement height;

u_R = velocity at the top of the roughness layer y_R .

- ✓ Nikora et al. (2000) put forward a concept that the large-scale turbulent eddies are in the order of the zero-plane displacement height d , where the mixing length l vanishes.

$$l = \kappa(y - d)$$

- ✓ The parameters κ , d , y_R and u_R/u_τ can be predicted from the fitting of $u(y)$ -profiles using experimental data.

- ✓ For low submergence, the wakes produced by the roughness layer reduce severely or eliminate the log-law and the outer layers, implying that κ does not hold.
- ✓ The evaluation of κ from the **log-law** loses its significance if the **roughness layer destroys** the **log-law** or the **outer layer**.
- ✓ Considering that the **roughness layer** extends up to $4k_s$ on the top of the roughness elements, the **non-existence of the log-law layer** is found up to $h \approx 4k_s$, ($h =$ flow depth).
- ✓ As a consequence, κ loses its implication and **cannot be evaluated**.

- ✓ This topic has attracted the attention not only of **hydraulicians** but also of **sedimentologists**.
- ✓ In the literature, there remain **many examples** of the **non-universality** of von Kármán's κ in flows over sediment beds:

For instance, von Kármán's κ behaves as a **variable** in flows with:

- (1) low relative submergence, $S_r (= h/k_s)$;
- (2) bed-load transport;
- (3) suspended-load transport.

Content

1. Introduction

2. Flows with Low Submergence

3. Flows with Bed-Load Transport

4. Flows with Suspended-Load Transport

5. Conclusions

Rand (1953):

- ✓ He was the first to conduct a flume experiment to study the effect of relative submergence on von Kármán's κ with artificial bottom roughness.
- ✓ He used 63 mm square bars spaced at 25.3 mm. He observed a value of $\kappa = 0.3$ for $S_r = 3.3$.

Bayazit (1976):

- ✓ He reported a gradually increasing κ (greater than 0.4) as the relative submergence S_r decreases below 2.5.
- ✓ He evaluated κ by fitting the velocity data to the logarithmic law having the position of virtual bed level at $0.33k_s$ below the roughness crest for all experiments.

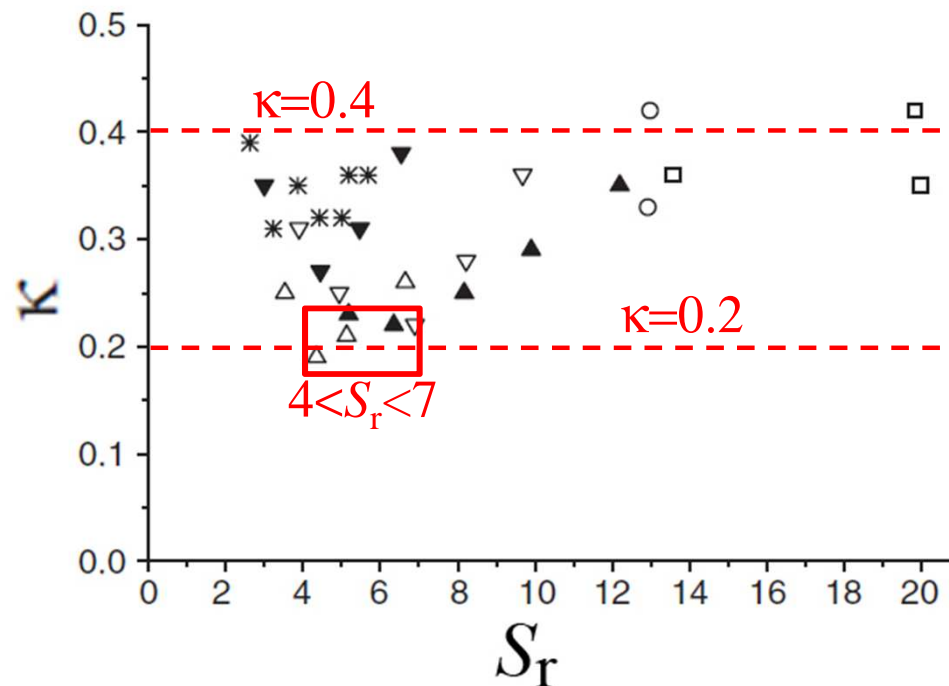
Pokrajac et al. (2006):

- ✓ Pokrajac et al. (2006) reanalyzed the experimental data of Bayazit (1976) with the roughness crest shear velocity instead of the bed shear velocity.
- ✓ Pokrajac et al. (2006) obtained modified values of κ . However, the values of κ remained greater than 0.41 for $S_r < 2.5$.

Dittrich and Koll (1997), Koll (2002) and Koll (2006):

- ✓ They showed that κ is non-universal and depends on both the irregularity of the surface and S_r .
- ✓ These parameters influence the formation and the evolution of the turbulence coherent structures, and in turn the velocity gradient.

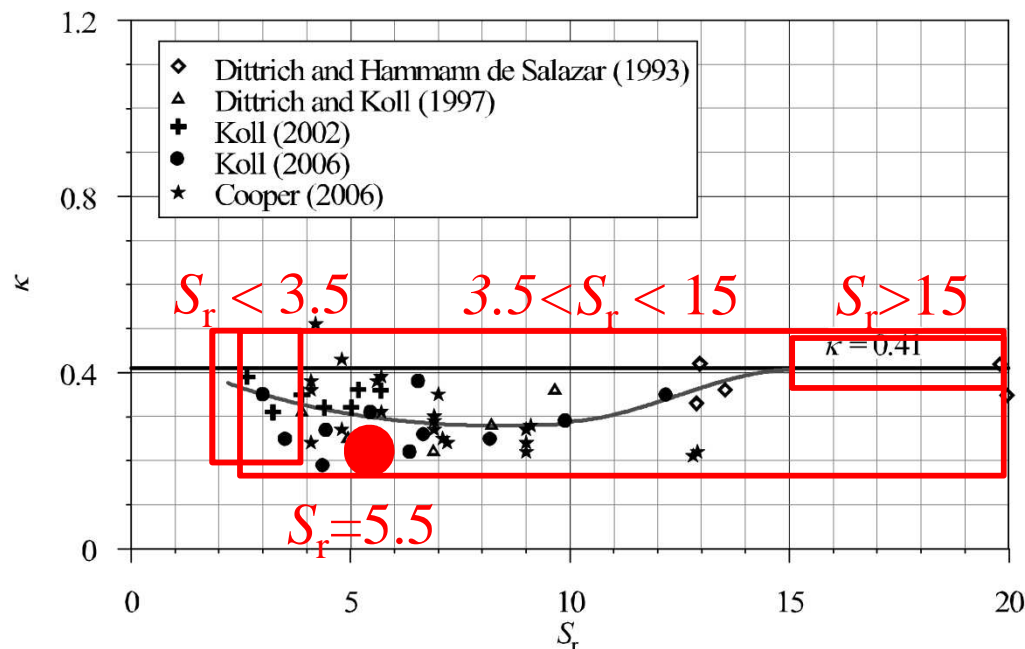
- ✓ Koll (2006) observed that κ approaches a value of 0.4 for regular surfaces and large values of S_r .
- ✓ It decreases significantly down to 0.2 if the bed roughness becomes large enough to decrease the flow depth relative to the roughness height.
- ✓ The value of κ reaches a minimum value within the range $4 < S_r < 7$ and increases again with a decrease in S_r .



Relation between κ and relative submergence S_r (Koll 2006)

Cooper (2006)

- ✓ also observed the values of κ lower than 0.41 for lower values of S_r ($4 < S_r < 13$).



Dependency of von Kármán's κ on the relative submergence S_r

- ✓ The figure shows that κ increases drastically for $S_r < 3.5$;
- ✓ It reaches the minimum value $\kappa = 0.25$ at $S_r = 5.5$;
- ✓ It becomes universal for $S_r > 15$;
- ✓ The curve κ versus S_r is sagging within $3.5 < S_r < 15$.

- ✓ Now the questions concerned with **the experiments** and the **estimations of the non-universality** of κ are:
 - 1) if Nikuradse's equivalent sand roughness k_s used for scaling h were the same as those obtained from the velocity profiles, and
 - 2) if roughness layer influences the value of κ .
- ✓ The **latter** is the **major decisive factor** towards the **non-universality** of κ ,
- ✓ the **former** provides the **scale up** to which κ is a **function** of S_r .

- ✓ Additionally, **hyporheic exchange** is often controlled by **subsurface advection**.
- ✓ It is driven by the **interaction of the fluid flow with sedimentary pore water**, as macro-rough beds are permeable.
- ✓ The nature and magnitude of the induced **hyporheic exchange** flow influences the main stream.
- ✓ This **influence** occurs by **changing the velocity and Reynolds shear stress** profiles and possibly κ as well.

Content

1. Introduction
2. Flows with Low Submergence
3. Flows with Bed-Load Transport
4. Flows with Suspended-Load Transport
5. Conclusions

Gust and Southard (1983)

- ✓ They observed a decrease in κ with an increase in bed-load transport.
- ✓ They suggested that, after a transitional regime related to the entrainment threshold of sediments, κ adjusted to a constant value of 0.32 ± 0.04 for all the bed-load experiments in which the transport rate varied by a factor 10.
- ✓ This implies that the uncertainties associated with the measured bed-load transport rate had a minimal effect on the result that κ is reduced by 25% from its universal value.

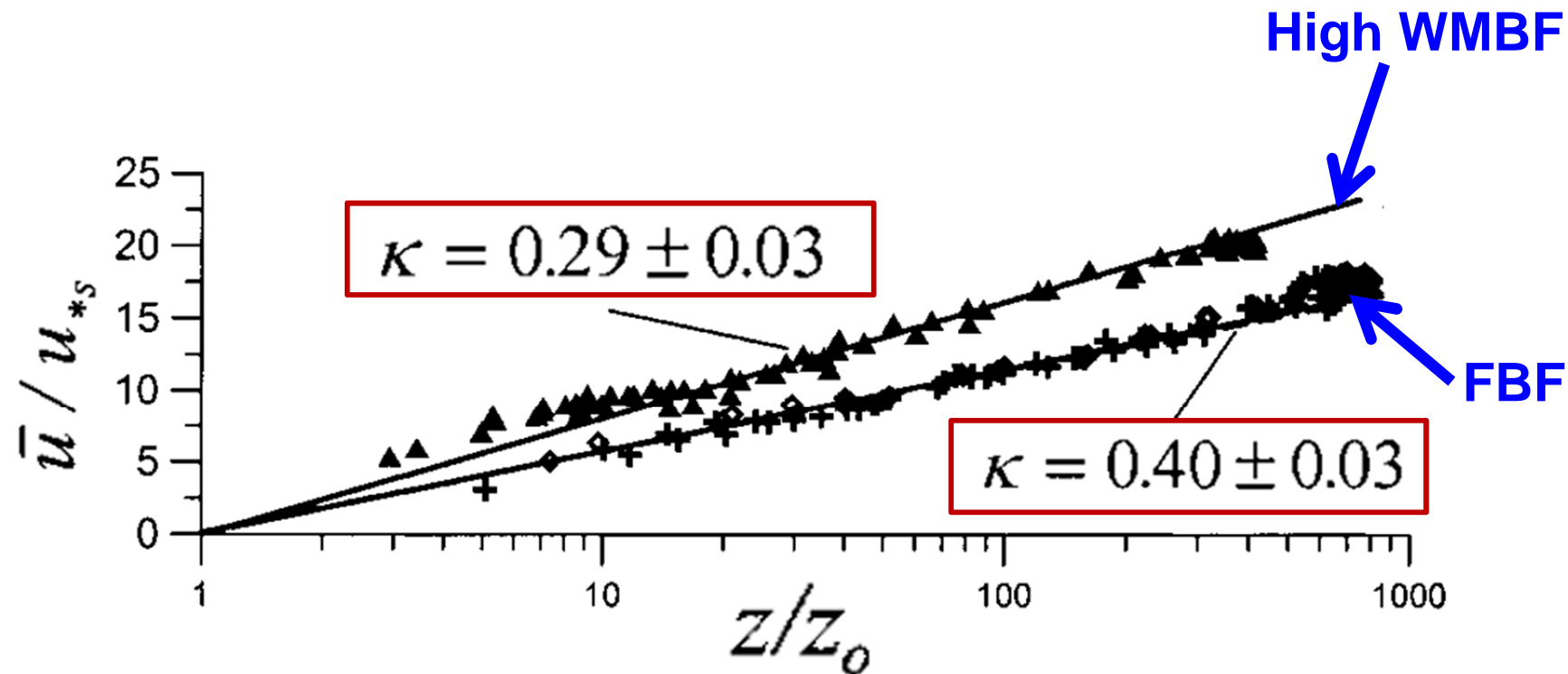
Best et al. (1997)

- ✓ They used a phase **Doppler anemometer** to differentiate the **characteristics** of the **fluid** from those of the **sediment particles**.
- ✓ They observed that the average value of κ was **0.385** in **presence of bed-load transport**.

Nikora and Goring (2000)

- ✓ They reported a study on the **characteristics of turbulent structure** of **high Reynolds number** in **quasi-two-dimensional flows**.
- ✓ They performed three sets of measurements taken with an **acoustic Doppler velocimeter** in an irrigation field canal for two bed conditions: (1) **fixed-bed flow (FBF)** and **weakly mobile-bed flow (WMBF)**.

- ✓ Nikora and Goring (2000) obtained $\kappa \approx 0.29$ for the **WMBF** being significantly smaller than $\kappa \approx 0.4$ for the FBF.



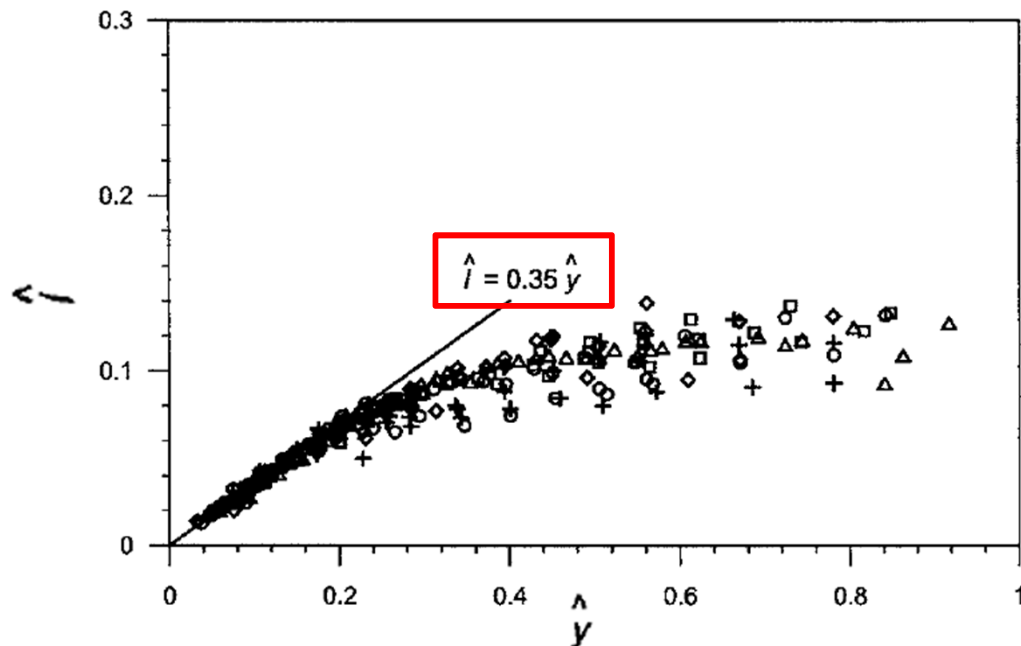
\bar{u} / u_{*s} : Local mean velocity/Friction velocity from turbulent stress distribution
 z / z_0 : Distance from the bed/Roughness length

- ✓ Bennett and Bridge (1995), Nikora and Goring (1999) and Gallagher et al. (1999) have also revealed an appreciable **decrease in κ** under **bed-load transport** conditions.
- ✓ Nikora and Goring (1999) imagined that the **reduction in κ** might reflect **the special turbulence characteristics** within a rather narrow range of the **Shields parameter** when the **bed shear stress** is approximately **equal to the critical shear stress**.
- ✓ In Nikora and Goring (2000), the **drag reduction** effects were expressed as **decreased values of κ** .

- ✓ The general concept is that the **drag reduction prevails** when the **spacing** between **turbulent bursting** events **increases** in comparison to the **spacing** in flows with **no sediment** (Tiederman et al. 1985).
- ✓ However, it is revealed that **κ reduces** when **spanwise** (lateral) **spacing** between **bursting events increases**, while **streamwise** spacing remains **unchanged** (Hetsroni et al. 1997).
- ✓ Nikora and Goring (2000) found that the **streamwise spacing** between **bursting events** was approximately **the same** for both **WMBF** and **FBF**, referring to an **increase** in **spanwise spacing** for **WMBF**.

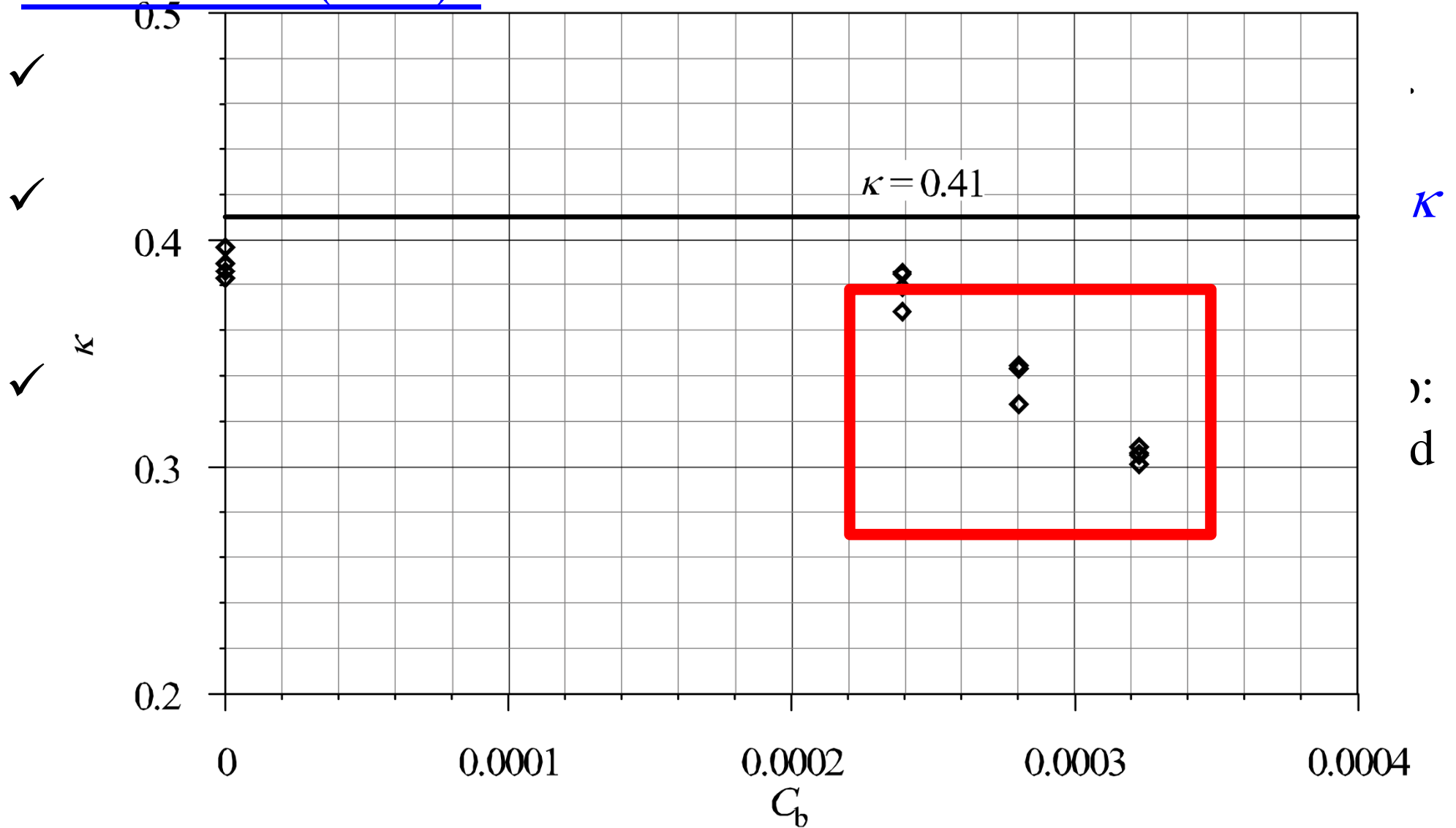
Dey and Raikar (2007):

- ✓ They reported the laboratory experimental results on the turbulent flow characteristics measured by an **acoustic Doppler velocimeter**.
- ✓ They observed that the **variation** of the **mixing-length** is considerably linear with the **elevation above the bed** within the **inner-layer** and obtained von Kármán's $\kappa=0.35$.



Mixing length as a function of flow depth

Gaudio et al. (2011):



- ✓ Further, Dey et al. (2012) fitted a **log-low** for **mobile-bed** flows to obtain $\kappa = 0.37$.

Summary of experimental data and results for κ in flows with bed-load transport

Source	d_{50} (mm)	S_x	B/h	R^*	F	Mode of transport / type of bed / sediment feeding	g_b (g/s/m)	κ
Gust and Southard (1983)	0.16	625	6	–	0.2 to 0.28	Bed-load / mobile bed / no feeding ($y/h \leq 0.2$)	0.0015 to 0.015	0.32 $\pm 12.5\%$
Best et al. (1997)	0.22	261	5.22	8.9	0.76 to 0.78	Bed-load / fixed bed / with feeding	9 to 22	0.385
Nikora and Goring (2000)	6.4	166	7.57 to 10.08	429	0.64 to 1.09	Bed-load / mobile bed / no feeding ($y/h \leq 0.2$)	13.8	0.29 $\pm 10.3\%$
Dey and Raikar (2007)	4.1 to 14.25	6.74 to 54.15	4.05 to 11.84	210 to 1573	0.17 to 0.38	Bed-load / mobile bed / near threshold ($y/h \leq 0.23$)	12.3 to 90	0.35 $\pm 0.86\%$
Gaudio et al. (2011)	1	52.5 to 67.81	2.9 to 3.6	101 to 120	0.98 to 1.01	Bed-load / fixed bed / with feeding ($y/h \leq 0.2$)	33.4 to 64.9	0.3 to 0.39 $\pm 10.7\%$
Dey et al. (2012)	0.95, 2.6, 4.1	29.2 to 158.0	6.3 to 34.2	63 to 508	0.55 to 0.77	Bed-load / fixed bed / with feeding ($y/h \leq 0.2$)	2 to 7	0.35 to 0.42

$F = U/(gh)^{0.5}$, U = mean flow velocity, and g = gravitational acceleration

Content

1. Introduction
2. Flows with Low Submergence
3. Flows with Bed-Load Transport
4. Flows with Suspended-Load Transport
5. Conclusions

- ✓ Also in flows with **suspended sediment** load von Kármán's κ has been a long **disputed parameter** in the data analysis of the **logarithmic velocity profile** fittings.
- ✓ During **1946-1961**, **pioneering research** was conducted by Vanoni (1946), Einstein and Chien (1955) and Elata and Ippen (1961) to examine the effect of suspended **sediment concentration** C_v on the **velocity profile**.
- ✓ One of the key findings was that κ **decreases** with an **increase in** C_v , although **Coleman (1981, 1986)** expressed a strong **dissatisfaction** on this issue.

- ✓ Coleman believed that Vanoni's finding was an artifact of the erroneous technique of evaluating κ that was generally accepted at the time when he did this work.
- ✓ According to Coleman, the experimental studies by Einstein and Chien (1955) were also unacceptable because the measurements were taken only over the lower 40% of the flow depth in the experimental channel that they used.
- ✓ Thus, it was difficult to determine the boundary layer thickness, the maximum velocity or any general information about the flow.
- ✓ The experiments by Elata and Ippen (1961) were done using virtually neutrally buoyant polystyrene particles to simulate the sediment suspension.

- ✓ The **velocity profiles** were presented in a velocity defect form; and the **apparent decrease of κ** was due to its **incorrect evaluation**.
- ✓ However, **Coleman (1981)** used the **wake-law** to study velocity profiles in **sediment-laden flows** and suggested that the **wake coefficient** rather than κ is **affected by the existence of sediment suspension**.
- ✓ The **wake-law** that describes the **logarithmic velocity** profile only **in the vicinity of the bed** has the same value of κ as that in **clear-water flow** ($\kappa = 0.41$).

- ✓ Lyn (1986) noted that the use of **pure log-law profile** leads to a **decreased** value of κ with C_v , while the use of log-wake-law leads to a universal one.
- ✓ Cioffi and Gallerano (1991) determined experimentally the **velocity** and **sediment concentration** profiles on a **mobile flat bed**.
- ✓ **without assessing κ in the inner region**, Cioffi and Gallerano (1991) only verified that the measured **velocity profiles** were reasonably interpolated with $\kappa = 0.4$ for $y/h < 0.15$.

- ✓ Cellino and Graf (1999) carried out an experimental study to investigate the influence of suspended sediments in the flow under non-capacity and capacity conditions.
- ✓ They accepted $\kappa = 0.4$ as a reasonable value for describing the velocity profiles through the velocity defect law.
- ✓ After Muste (2002), a constant value for $\kappa = 0.41$ would be appropriate for the sediment-laden flows involving only small C_v (smaller than 0.05).
- ✓ The aforementioned discussion is given on the investigators who opposed the changed value of κ from its universal value.

- ✓ On the other hand, other investigators believed that κ is non-universal in sediment-laden flows.
- ✓ Many other researchers during 1960s and 1970s also reported that κ diminishes as C_v increases (Vanoni and Nomicos 1960; Hino 1963; Paintal and Garde 1964; Bohlen 1969; Ippen 1971).
- ✓ Such decrement is primarily due to increasing gradient of the velocity profiles in presence of sediment suspension.

Nouh (1989):

- ✓ He reported that the variation of κ with the average concentration C_{av} of suspended sediments in straight open-channel flows depends on the flow Reynolds number, $Re = 4Uh/\nu$.
- ✓ As C_{av} increases, κ decreases in flows with $Re < 7 \times 10^5$, but increases in flows with $Re > 7 \times 10^5$.
- ✓ The variation of κ with C_{av} is insignificant in flows with Re equal to about the critical value $Re_c = 7 \times 10^5$.
- ✓ He hypothesized that the turbulence level close to the boundary decreases for $Re < Re_c$ and increases for $Re > Re_c$, as C_{av} increases.

- ✓ For $Re < Re_c$, C_{av} affects the turbulence level more than Re , and vice versa for $Re > Re_c$. These two effects are balanced for $Re \approx Re_c$.
- ✓ Nouh (1989) also observed that the increased values of κ in flows with fine suspended sediments are larger than those in flows with relatively coarse suspended sediments, and also those in flows with high C_{av} are larger than those in flows with low C_{av} .
- ✓ In fact, he explained that, for a given C_{av} , coarse suspended sediments produce a larger reduction in turbulence level than fine suspended sediments.
- ✓ In a clear-water flow, universal value of κ is invariant of Re (for $4 \times 10^5 < Re < 2 \times 10^6$).

Wang and Qian (1992):

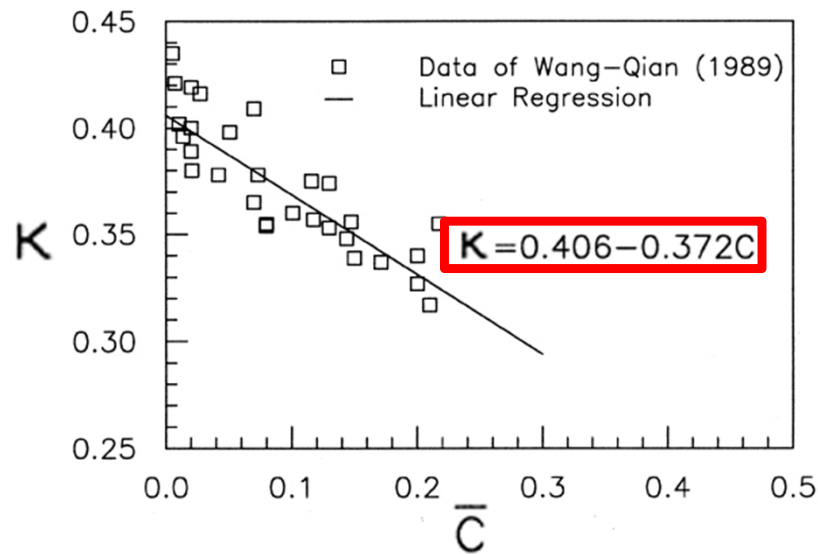
- ✓ Wang and Qian (1992) showed that, in the lower-flow region, the values of κ in sediment-laden flows are less than those in clear-water flows.

Guo and Julien (2001):

- ✓ They argued that the reduction of κ in sediment-laden flows is governed by C_{av} and the mass density gradient given by the Richardson number.

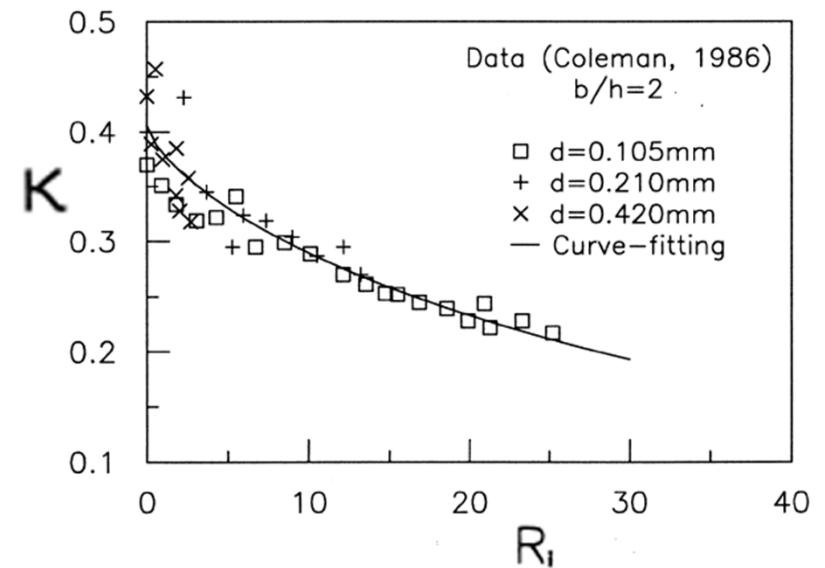
Analysis by Guo and Julien (2001):

Average concentration effect on von Kármán's k



\bar{C} : Time-averaged volumetric sediment concentration

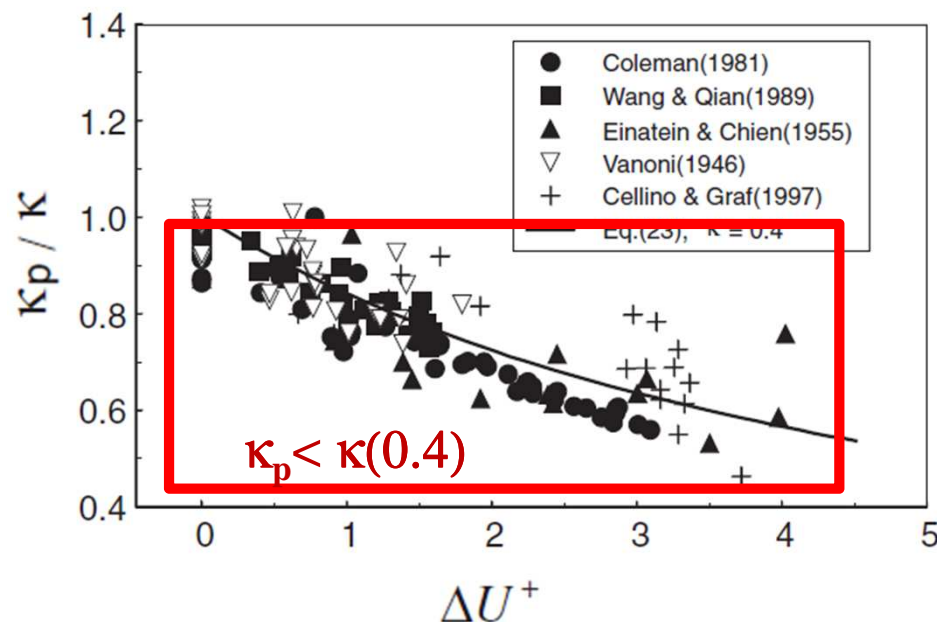
Density gradient effect on von Kármán's k



b/h : Channel aspect ratio
 d : Particle diameter
 R_i : Richardson number

Wang et al. (2001):

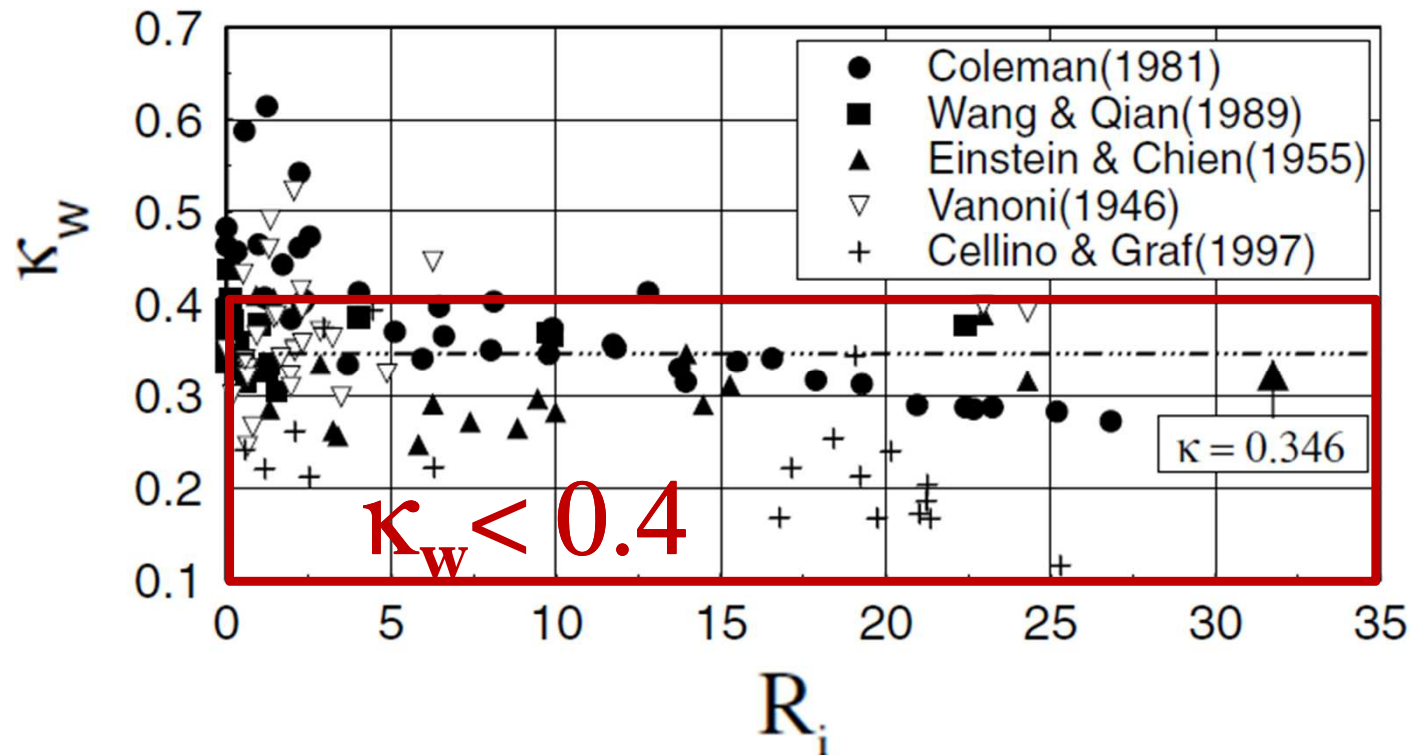
- ✓ In presence of **sediment suspension**, they modified von Kármán's κ_p for the **log-law** and κ_w for the **wake-law** separately.
- ✓ For the **log-law**, they proposed $\kappa_p = 2.08\kappa/(\Delta U^+ \kappa + 2.08)$, where ΔU^+ is $(u_c - u_s)/u_\tau$ and u_c and u_s are the velocities at $y/h = 0.05$ in **clear-water** and **sediment-laden** flows, respectively.



The relationship of κ_p and ΔU^+

Wang et al. (2001):

- ✓ For the wake-law, they obtained the average value $\kappa_w=0.346$ with high data scattering.



Relationship of κ_w and R_i

Wang et al. (2001):

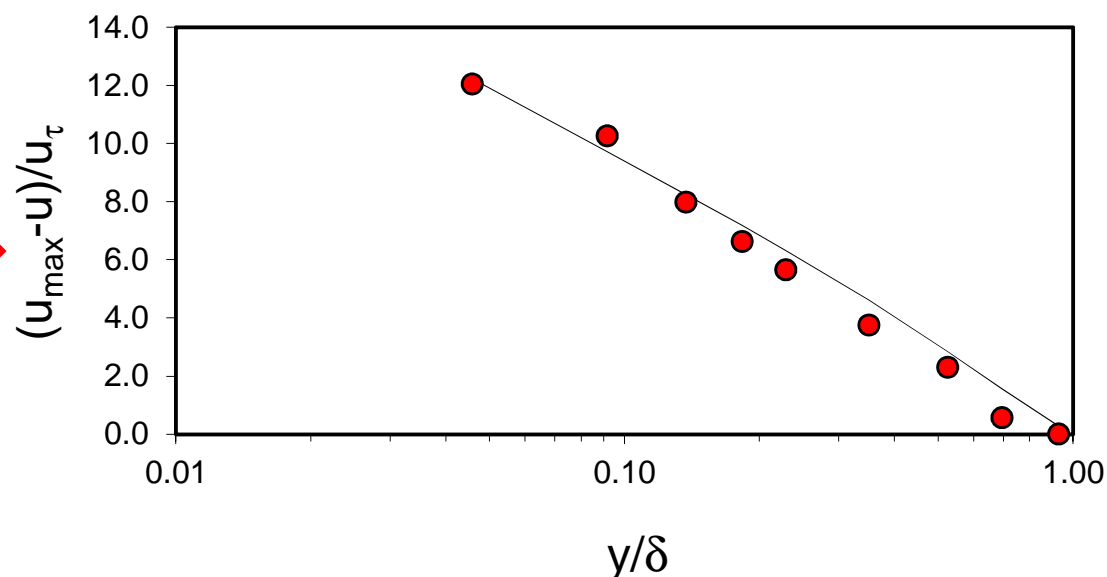
- ✓ They conducted simultaneous measurements of both the suspended particles and water in particle-laden flows.
- ✓ They used a discriminator particle-tracking velocimeter, observing that κ decreases with an increase in C_{av} .

Reanalyzing Coleman's data:

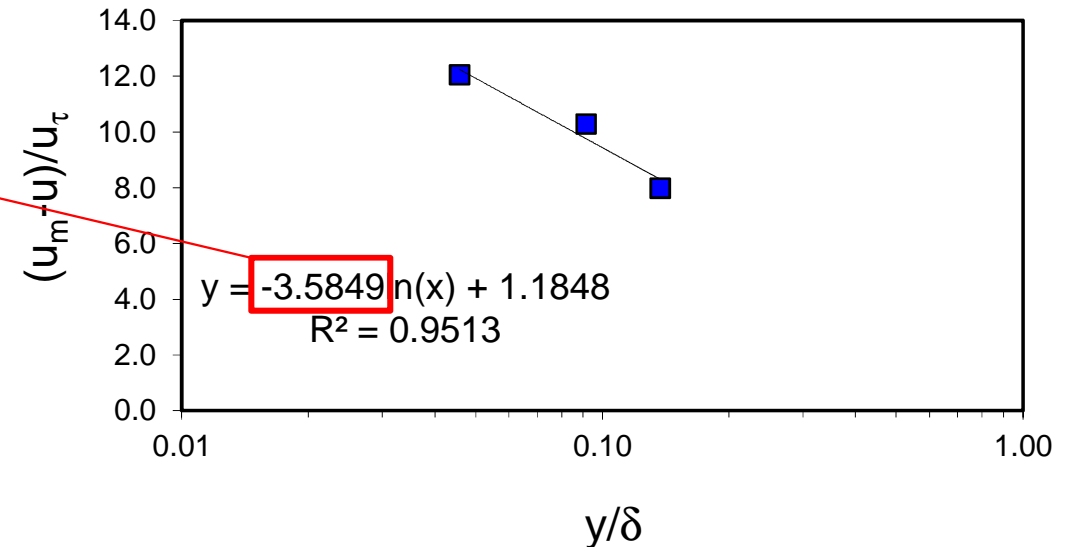
- ✓ In fact, the investigators supporting the universality of κ also in sediment-laden flows referred to Coleman's (1981, 1986) data and analysis.
- ✓ Owing to the available evidence that κ varies in the presence of suspended sediments, it seems to be necessary to reanalyze Coleman's dataset in order to verify the proposed constancy of κ .

- ✓ By adopting the **velocity defect law** as in the original work, the dimensionless data $(u_{\max} - u)/u_{\tau}$ were plotted as a function of y/δ in a semi-log graph;
- ✓ δ , the **boundary layer thickness**, is the distance above the bed at which the **time-averaged** streamwise velocity reaches its maximum value u_{\max} .

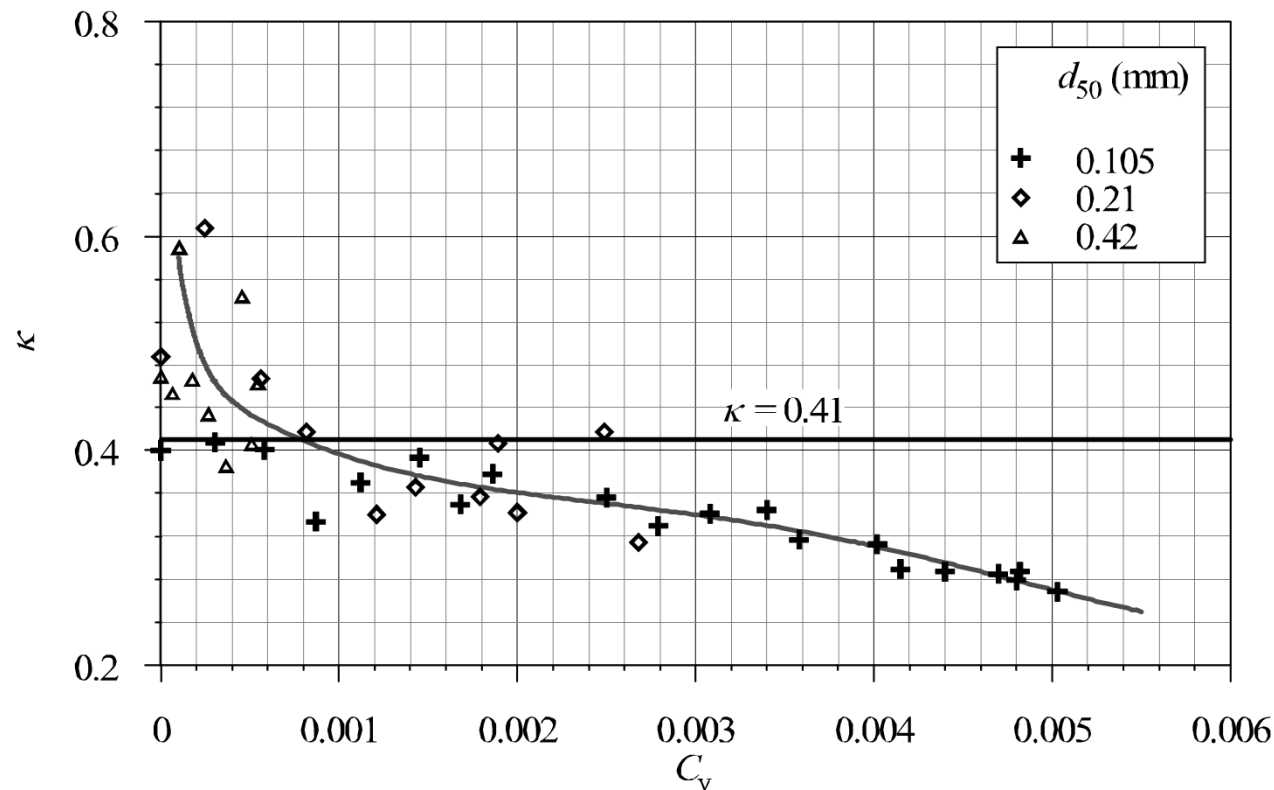
Example: Coleman's Run 19



$$\begin{aligned}
 -\kappa^{-1} &= -3.5849 \\
 \kappa &= 0.279
 \end{aligned}$$



- ✓ The slopes $-\kappa^{-1}$ were evaluated from **the regression analysis** fitting straight lines within the log-law layer in the inner flow region ($y \leq 0.15\delta$).
- ✓ We obtained values of κ different from those given in **Coleman (1981, 1986)**.



Variation of κ with C_v obtained by reanalyzing Coleman's data (1981, 1986)

- ✓ The mean trend (solid gray line) shows progressively diminishing values of κ as C_v increases.
- ✓ In general, for $C_v > 0.0008$, the values of κ are less than its clear-water value 0.41.

Content

1. Introduction
2. Flows with Low Submergence
3. Flows with Bed-Load Transport
4. Flows with Suspended-Load Transport
5. Conclusions

- ✓ Reasons of the variability of κ are recognized as follows:
1. Sediment particles during the **bed-load transport** interact with both the **flow that accelerates** them and the **rough bed that decelerates** them (Gyr and Schmid 1997).
 2. **Collisions** cause **momentum extraction** from the mean flow in the near-bed region, resulting in **reduction of local streamwise velocities** and **increase of velocity gradients** (Owen 1964; Smith and McLean 1977).
 3. In **low submergence**, κ was found to depend on the **irregularity of the surface** and the **relative submergence**, which influence the formation and the expansion of the coherent turbulent flow structures, and thus the **velocity gradient**.

- ✓ In case of **bed-load transport**, it was not even possible to provide any solid relationship between κ and the bed-load transport rate, although Gaudio et al. (2011) provided an empirical relationship expressing κ as a function of C_b and R_b/d_{50} .
- ✓ The von Kármán κ is given by l/y , where l is the Prandtl mixing length, $[-\overline{u'v'}/(du/dy)^2]^{0.5}$, and u' and v' are the fluctuations of the streamwise and vertical velocity components, u and v , respectively.
- ✓ Thus, κ is **dependent on** the **Reynolds shear stress** $-\overline{u'v'}$ relative to the **mass density** of fluid and on the **velocity gradient** du/dy .

- ✓ Hence, the variable values of κ depend on the prevailing effect between turbulence intensity and velocity gradient.
- ✓ For instance, if turbulence intensity increases and/or velocity gradient decreases, then κ increases and vice versa.
- ✓ In order to assess the non-universality of κ , future researchers have to keep an eye on the variations of these two parameters with vertical distance within the logarithmic law layer.

Thank you for your kind attention

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Thank you for your kind attention